Implications of the Rupture Process from the Displacement Distribution of Strong Ground Motions Recorded during the 21 September 1999 Chi-Chi, Taiwan Earthquake

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ABSTRACT

The 21 September 1999 Chi-Chi earthquake (M, =7.3) was the largest inland quake occurred in Taiwan in the past 100 years. According to direct field observations, a significant thrust rupture was occurred along the active Shuangtung fault and Chelungpu fault. From the integrated displacement derived from free-field strong motions around the source area, horizontal displacements over 9 m and vertical offsets reaching 3.4 m are obtained at the station TCU068 located nearby the northern section of Chelungpu fault. Based on the field observations of slip distribution over the entire 77-km surface break and the integrated displacement waveforms, we have arrived at the following findings: An obvious small offset at about 3 sec prior to the principal offset is observed at two sites close to the southern part of Shuangtung fault. This small offset is not shown on the other stations. We propose that a relatively small event, probably with a magnitude of about 6.0, occurred about 3 sec before the major rupture, that is located at the Shuangtung fault about 8 km east of the Chelungpu fault. The induced northward propagating rupture along the Chelungpu fault has reached a maximum permanent offset over 9 m at the northern section of fault. This northern section of the Chelungpu fault was identified as one of the heaviest damaged areas.

> (Key words: Chi-Chi earthquake, Fault rupture, Permanent displacement field)

1. INTRODUCTION

A major earthquake ($M_L=7.3$) has occurred near the town Chi-Chi in the Nantou county, Taiwan, at 1:47 a.m. (local time), September 21, 1999. The focal depth is 8 km. The death toll until the end of September has exceeded 2100 and is mounting during the first week after the main shock. This earthquake is the largest inland event in Taiwan in this Century and has

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provided a wealth of digital data for seismology and earthquake engineering because a successful implementation of the six-year Taiwan Strong Motion Instrumentation Program (TSMIP) completed two years ago. It is believed that this big earthquake is intimately related to the recurrent movement of the Shuangtung fault and the Chelungpu fault. The Chelungpu fault is a thrust fault, that has an extension of about 77 km from Minchien, Nantou in the south to Fengyuan, Taichung to the north, with an almost N-S strike (Chang et al., 1998). According to the direct field investigation, the surface break shows large vertical as well as horizontal displacements along the fault line as a consequence of the Chi-Chi earthquake.

Over thirty free-field strong-motion stations equipped with an A-900 (Teledyne Geotech) or IDS-3602 (Terra Technology) accelerometer are deployed close to the Chelungpu fault by the Central Weather Bureau (CWB) under the TSMIP project, which lasts from 1991 to 1997. From those retrieved near-field accelerograms of this Chi-Chi event, the significant permanent offsets are easily derived by using a simple integration process. These displacement results give the reason why the heaviest damage occurs along the northern section of the Chelungpu fault zone, in complete agreement at the on-site damage investigation provided by many researchers. For such a devastating event, the strong-motion instruments deployed nearby the fault can record the dynamic behavior associated with the rupturing process. This paper shows the displacement field on the both sides along the two faults through a simple integration of the accelerogram data. This result provides direct evidence to geologists and seismologists in the interpretation of faulting process. The rupture characteristics of this seismic source, as it relates to the regional tectonics, can also be reconstructed with this integrated displacement field as an important constraint.

2. DATA PROCESSING

For the TSMIP network, each station (either an A-900 or an IDS-3602) consists of a three-component force-balanced accelerometer recorded at a 16-bit resolution and a 96-db dynamic range. The full scale of recording system is $\pm 2g$. The instrument response of acceleration is almost flat from DC to 50 Hz. Basically, the basic specifications of the both instrument types are the same, except the sampling rate. The A-900 is sampled at 200 sps, and the IDS-3602 at 250 sps. It is necessary to do some pre-processing, like decoding, time-checking, and baseline correction, on the raw data.

During the first five days after the main shock, some accelerograms recorded in the central Taiwan region are retrieved and are immediately sent to CWB for data processing. To obtain the displacement field, the accelerograms are integrated twice in the frequency domain. Suppose the acceleration is A(t), then it can be transformed into frequency domain as A(ω) by using the Fourier transform method, where t and ω represent the time and the angular frequency, respectively. By the integration theorem, the displacement D(t) can be derived from the inverse Fourier transform of D(ω), which is equal to $-A(\omega)/\omega^2$. Obviously, passing over to the frequency domain, the operations are straight-forward. Since all the data used in this paper have very high signal-to-noise ratio, there is no need to apply additional filtering during the integration process. Inevitably, the displacement seismograms still have some long-period components caused by the instrumental drift during the time of strong shaking or the digitization problem in the part prior to the event signal. Although the anomaly produces only a small acceleration offset, it can have a noticeable effect on the double integral of the acceleration data (Iwan et al., 1985). In general, the long-period (200 to 400 sec) trend is obvious and can be easily traced in the integrated displacement waveform, since the record length is over 60 sec that is much longer than the displacement signal. Therefore, we use a sinusoid function to fit the trend and subtract it from the integrated time history to obtain the displacement waveform for measuring the permanent offset.

3. RESULTS

Figure 1 is the accelerogram recorded at station TCU068 located at about 8 km east of Fengyuan. The peak ground accelerations (PGA) in the U-D, N-S, and E-W components are 519 gals, 362 gals, and 502 gals, respectively. The integrated displacements that the long trend is removed are shown in Fig. 2. Obvious ramp-type slip function appears on the displacement waveforms. In general, it is not difficult to estimate the permanent offsets directly from the waveform for such a case. The estimations show that the first major movement is about 451 cm to up, 776 cm to north, and 820 cm to west with a rise time about 4~5 seconds. Followed with the first movement, the elastic rebound with 106 cm, 266 cm, and 55 cm, respectively, is occurred and lasts about 3 seconds. Therefore, the permanent offset is 345 cm to up, 510 cm to north, and 765 cm to west. This clear observation of the slip function at this large amplitude is perhaps the first time in seismological history. Another station TCU052, located 10 km south of TCU068, has also recorded extremely large offset, about 340 cm to up, 550 cm to north, and 345 cm to west (Fig. 3). These two stations are both located at northern part of the hanging wall of the Chelungpu thrust fault. On the other hand, one of the stations near southern part of the fault, namely TCU129, has recorded 983 gals (over 1 g) in the E-W component of acceleration, which is about two times greater than the PGA recorded at the above two stations. But the horizontal offset on this site is only 50 cm to south and 110 cm to east. No obvious offset is estimated in the vertical component. This phenomenon indicates that the slip along the fault rupture is not necessary proportional to the generated acceleration. The effects of the rupture propagation should play an important role in this case. A forthcoming paper will further clarify that point.

The derived offsets at twenty-two CWB's free-field strong-motion stations near the Chelungpu fault and the Shuangtung fault are given in Table 1 and the displacement distribution around the two faults is shown in Fig. 4. The two orthogonal red arrows at each site mean the horizontal N-S and E-W displacements, while the solid and open blue arrows represent upward and downward movements, respectively. The three component displacements for all stations are plotted in the same scale indicated in the figure legend. It is clearly seen that the offsets on the eastern side of Chelungpu thrust fault, the hanging wall, have systematically larger northwestern displacements, which are opposite to the movements of the western side, the footwall. With the well-known southeast-northwestern regional tectonic compression operative in Taiwan, the offset distribution gives a consistent picture and has explained the major



Fig. 1. Accelerogram recorded at the free-field station TCU068.

thrust faulting along the Chelungpu fault during the earthquake. Field geological survey of the surface break also gives the ground truth of this integrated displacement field. The offset decreases to only several centimeters on the sites only 10 km away from the fault on the footwall side.

For a thrust fault, the ground motion is not expected to be the same on both sides of the fault. Based on simple geometry, it is expected that sites located on the hanging wall will have larger ground motions than sites at the same rupture distance located on the footwall because the hanging-wall sites are closer to a larger area of the source than the footwall sites. However, the two stations, TCU068 and TCU052, with the extreme large offsets should be located



Fig. 2. Integrated displacement waveforms derived from the accelerogram recorded at the free-field station TCU068. The permanent offset for each component is measured directly from the waveform.

on the hanging wall side and are very close to the rupture outcrop of Chelungpu fault. Examining the displacement distribution in the area, the N-S offset seems increase from south to north of the fault zone. This phenomenon implies that the pure thrust rupture propagates and changes from pure thrusting to oblique thrusting with strike-slip component from south toward north, with increasing rupture slip. For the vertical permanent offset distribution, only a small downward slip on the order of about 10 cm are observed on the footwall sites, while larger uplift can be derived on the hanging-wall sites. The pattern agrees well with the static



Fig. 3. Integrated displacement waveforms derived from the accelerogram recorded at the free-field station TCU052.

model of the thrust faulting. Some downward slips are also measured at stations located southeast of Shuangtung fault. It is due to the pulling of up-thrusting material that causes the sinking of soft surface sediments under the extension.

4. DISCUSSION AND CONCLUSION

The location of the main shock epicenter is at 120.82°E, 23.86°N, where is about 2 km the eastern side of the south tip of Shuangtung fault. Figure 5 represents the focal mechanism

Station	Longitude	Latitude	Final displacement (cm)			PGA (gals)		
Code	°Е	°N [U-D	N-S	E-W	U-D	N-S	E-W
TCU068	120.766	24.278	345	510	-765	519.4	361.9	501.6
TCU102	120.721	24.249	-13	-58	93	173.3	169.0	298.4
TCU053	120.669	24.194		-68	75	120.9	132.1	224.8
TCU052	120.740	24.198	340	550	-345	194.0	438.7	348.7
TCU049	120.690	24.180	—	-55	75	178.0	241.9	273.4
TCU054	120.675	24.162	—	-90	60	133.0	190.4	143.1
TCU082	120.676	24.148	·	-50	92	129.3	182.5	221.1
CHY025	120.514	23.780	.—	-8	30	169.7	152.0	158.6
CHY101	120.562	23.686		-30	-35	162.2	390.1	332.7
TCU067	120.720	24.092	-13	-40	140	230.6	312.7	488.9
TCU065	120.691	24.059	-12	-30	125	257.8	563.2	774.4
CHY024	120.607	23.758	-10	-7	48	141.4	162.2	276.3
TCU072	120.849	24.041	145	225	-190	274.7	370.5	465.3
TCU122	120.610	23.813	-4	-16	75	236.0	255.6	207.4
TCU075	120.678	23.984	-12	-16	120	223.9	257.3	325.3
TCU074	120.962	23.961	-58	132	-190	270.2	368.4	585.9
TCU076	120.676	23.908	-5	-42	75	275.4	420.0	340.1
TCU129	120.684	23.878	—	-50	110	335.0	610.8	983.0
TCU089	120.857	23.904	-10	130	-175	190.2	225.3	347.8
TCU078	120.846	23.812	-17	65	-110	171.0	302.5	439.7
TCU071	120.788	23.986	255	315	-250	415.5	639.0	517.8
TCU120	120.613	23.980	-2	-22	70	166.5	193.5	223.0

Table 1. The Final Displacements and PGA at Free-field Strong-motion Stations.

The symbol '— ' in U-D displacement indicates that the value is about zero or is difficult to estimate.

** The negative values of final displacement in U-D, N-S, and E-W component represent downward, southward, and westward movement, respectively.

determined from the polarity of P waves recorded by the CWB short-period seismic network and some strong-motion stations. The solution is with the strike N20°E, dip 30°, and rake of 85°. It is presumably the rupture plane at its nucleation stage. Figure 6 shows the displacement waveforms recorded at the stations TCU078 and TCU089 located at the eastern side of Shuangtung fault. These two stations are the closest sites from the epicenter. The integrated displacement waveforms show a smaller slip observed at about 3 sec prior to the principal slip pulse. This smaller slip is not observed at the other stations. This phenomenon indicates that a relatively small event, probably with a magnitude of about 6.0, occurred before the major rupture and is estimated at about 8 km east of Chelungpu fault. Using the final displacement





distribution observed at the source area, we propose that the Chi-Chi earthquake initiate at the Shuangtung fault and then the major rupture at Chelungpu fault was triggered in about 3 seconds. The northward propagating rupture along the Chelungpu fault has reached to a maxi-

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mum offset at the northern section of fault where the free-field strong-motion stations TCU068 and TCU052 are located. The estimation of the slip rate on the northern section of Chelungpu fault is in the range from 100 to 200 cm/sec from these two displacement records.

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