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Area Balancing Method as a Useful Tool to Estimate the Depth of the Chi-Chi Earthquake Rupture at the Fengyuan Well

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ABSTRACT

The 1999 Chi-Chi earthquake, $M_w = 7.6$, was associated with a near 100-km-long surface rupture along which the horizontal slip and vertical displacement reached to 9 m and 4 m, respectively, according to coseismic GPS slip vectors. After the Chi-Chi earthquake, drilling projects were conducted to clarify the characteristics of the earthquake rupture in depth. At the Fengyuan site, inclined-well drilling was performed at an angle of 50 degrees westward to a drilling depth of 455.3 m in the northern segment of the Chi-Chi earthquake rupture. Two possible earthquake rupture zones were observed at drilling depths of 225 m and 330 m according to the core structures and seismic profiling. We apply an area balancing method to determine the depth of the earthquake rupture in the cores; and we suggest that the 330-m drilling depth is the most likely candidate for the earthquake rupture.

(Key words: Chelungpu fault, Area balancing method, Chi-Chi earthquake, Fengyuan well)

1. INTRODUCTION

The 21 September, 1999 Chi-Chi earthquake ($M_w = 7.6$), a consequence of the on-going collision between the Philippine Sea plate and Eurasian plate (Fig. 1), produced a surface rupture about 100 km long in western Taiwan. At the northern segment near the Shihkang area, the rupture has a throw of up to 10 m, the largest coseismic displacement ever docu-

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Fig. 1. (a) Horizontal and vertical displacements of the GPS at the northern segment of the Chi-Chi earthquake. (b) In the northern segment of the Chi-Chi earthquake, the rupture of Chi-Chi earthquake shows a rampflat structure (Lee et al. 2005). The vertical displacement is larger than 3 m on the ramp plane and near 1 m on flat plane.

mented on a reverse fault (Chen et al. 2001a; Yu et al. 2001) (Fig. 1). The Chi-Chi earthquake was caused by a slip along the Chelungpu fault, a major thrust fault in the frontal fold-and-thrust belt of western Taiwan.

After the Chi-Chi earthquake, many drilling projects and trenching studies have been conducted to analyze the characteristics of the rupture of the Chi-Chi earthquake (Chen et al. 2004). For example, drilling projects have been conducted at both the Fengyuan and Nantou areas (Wang et al. 2000, 2002). At the Nantou site, the drilling well successfully penetrated the fault at around a depth of 177 m (Tanaka et al. 2002). In this well, the position of rupture in the core is easily recognized by different strata in the footwall and hanging wall. At the Fengyuan site, the drilling was performed as an inclined well at an angle of 50 degrees westward (Fig. 6). It appears to penetrate possible earthquake rupture zones at drilling depths of 225 m and 330 m (Tanaka et al. 2002). According to shallow seismic reflection studies and core structures, the 225-m drilling depth was considered the position of the Chi-Chi earthquake slip plane (Tanaka et al. 2002). However, some researchers consider the 330-m drilling depth to be the actual position of the slip plane rupture based on the fracture density distribution of the fault zone (Heermance et al. 2003). This ambiguity results from the facts that both the hanging wall and footwall of the Chi-Chi earthquake rupture have similar strata at the Fengyuan site, and that the existence of many fractures increases the difficulty in identifying the Chi-Chi earthquake slip plane. Identifying the position of the slip surface in the drilling core has become a key issue; and this represents one of the main targets of this study. This research utilizes the area balancing method to better identify the depth of the slip plane. As a result, we suggest that the position of the earthquake slip plane at the Fengyuan well is at the 330-m drilling depth.

2. GEOLOGICAL BACKGROUND

The Chelungpu fault is a major thrust fault in the western foothills of Taiwan where the Pliocene to Miocene strata thrust over on the Pleistocene Toukoshan Formation with a tenkilometer net separation (Fig. 2) (Chang 1971). In a strict sense, only the middle to southern segments of the Chi-Chi earthquake rupture follows the surface trace of the Chelungpu fault. At the northern segment, the rupture does not follow the Chelungpu fault, but rather a new fault that is a bedding slip along the Chinshui Shale (Fig. 2) (Chen et al. 2001b; Lee et al. 2000; Lee et al. 2003a; Lee et al. 2005).

At the northern segment, to the north of Wufeng, the subsurface structure of the Chi-Chi earthquake fault shows a ramp-flat structure with vertical displacement of more than 3 m on the ramp plane and less than 1 m vertical displacement on the flat plane (Fig. 1) (Lee et al. 2005). The ramp plane is located from the western side of the Shihweichiao syncline axis to the surface rupture. The strata strikes roughly N-S with 40° - 70° dipping to east near the surface rupture and progressively decreasing to 30° toward the east on the ramp plane (CPC 1974, 1982; Lee 2000b; Lee et al. 2005) (Fig. 2). Using offset of infrastructures (e.g., city streets or concrete fences), Lee et al. (2003a) measured nearly 100 sets of slip vectors along the surface rupture that show the slip direction is in NW from Fengyuan to Wufeng, which is consistent with the GPS data.



Fig. 2. Geological map (CPC 1974, 1982) and drilling sites at Fengyuan and Takeng area.

3. THE AREA BALANCING METHOD

Area balancing method is a classic technique for calculating the depth of a detachment such that the increasing area in front of a measured point will be equal to the depth of detachment multiplied by the known shortening of strata (Dahlstrom 1969). We apply the same concept to estimate the depth of the earthquake slip plane with significant coseismic slip amount.

We assume, in a given cross section parallel to the regional coseismic slip trend of the thrust block, the area of the section is conserved that means there is no volume loss and no density change. Accordingly, at a given drilling site, the amount of the area horizontally moving along the slip surface $(A \times C)$ is equal to the area $(D \times V)$ gained due to uplift next to the site (Lee et al. 2003b; Lee et al. 2005) (Fig. 3). Assuming that the slip of the thrust block is constant from the surface to the depth under shallow depth condition, the total area transferring along the detachment surface can be approximated by the horizontal slip of the thrust block multiplied by the thickness of the block (i.e., the depth of the detachment surface). That means if we know the horizontal displacement, distance from measuring point to the surface rupture, and vertical displacement, we can easily calculate the depth of the earthquake rupture (Fig. 3). Using the GPS data, Johnson and Segall (2004) derived geometry of the Chelungpu fault according the dislocation model. Figure 4b shows the horizontal displacement distribution from surface to slip plane according to the results of Johnson and Segall (Johnson and Segall 2004). The horizontal displacement is not the same from surface to slip plane in depth but does not change a lot in the near surface rupture. This supports, to some extend, the assumption that the horizontal displacement is the same from surface to slip plane. We also tested the area balancing method at the Neiwan area (Lee et al. 2005). According to seismic profile and well data, the depth of the Chinshui Shale is about 600 m at the Neiwan area (Fig. 5b). The Chinshui Shale is the slip plane of the Chi-Chi earthquake. The Neiwan area shows a coseismic folding



Fig. 3. Schematic map of balancing-cross section method. The increased area is equal to horizontal displacement multiplying with the depth of the earth-quake rupture.



Fig. 4. Horizontal displacement distribution from surface to fault plane. (a) The geometry and slip distribution of the Chi-Chi earthquake is according to Johnson and Segall (2004). The strike of the AA' is parallel to the displacement direction of the northern segment of the Chi-Chi earthquake.
(b) The horizontal displacement amount is similar from surface to fault plane, and it progressively decreases away from the surface rupture. We constructed this profile by Poly3D (Thomas 1993).

structure during the Chi-Chi earthquake. The increased area of the folding structure is about 5970 m^2 and the horizontal slip is 10.4 m; therefore we can obtain the depth of the detachment to be about 573 m, which is roughly consistent with the depth indicated from well and seismic profiling (Lee et al. 2005) (Fig. 5). This result indicates the usefulness of the area balancing method.



Fig. 5. Depth of the slip plane of the Chi-Chi earthquake under the Neiwan area.
(a) Deformation profile in the Neiwan area. The increased area is 5970 m² and the calculated depth of the Chi-Chi earthquake rupture is 573 m.
(b) The seismic profile around the Neiwan area. The location of the profile shows in Fig. 1. The slip plane was along the Chinshui Shale, which is at about 600 m in depth and this depth is similar to the calculated value.

4. VERTICAL DISPLACEMENT AROUND THE DRILLING SITE

While we apply the above area balancing method to estimate the depth of the earthquake slip plane, the vertical displacement is an important factor affecting the accuracy of the measurement. The scarp is about 3 - 4 m high around the Fengyuan area according to direct field measurements (CGS 1999; Lee et al. 2003b) and 1 : 1000 scale topographic maps along the surface rupture (Fig. 6) (NCREE 2000).

The northern segment of the Chi-Chi earthquake rupture shows a ramp-flat structure and



Fig. 6. The vertical displacement is 3 - 4 m high along the surface ruptures of the Ch-Chi earthquake. The bottom of figure shows the measuring results of the balancing method and fault geometry. The dotted line is the fault geometry according to 3.3-m vertical displacement. The gray area shows the possible position of the fault plane.

the Shihweichiao syncline is located on the kink axis of the ramp-flat structure. According to cosesimic GPS data, the vertical displacements are 3 - 4 m and ~ 1 m high on the ramp and flat structure, respectively (Yu et al. 2001, Yang et al. 2000, Lee et al. 2003a, b; Lee et al. 2005) (Fig. 1). The Fengyuan drilling sites are located on the ramp structure. Therefore, we assumed that the regional vertical displacement from the drilling site to the surface rupture is about 3 - 4 m, although we have no exact vertical displacement data.

5. DEPTH OF THE EARTHQUAKE RUPTURE AT FENGYUAN WELL

The Fengyuan site is about 2.1 km ENE from Fengyuan city and is about 500 m to the east of the surface rupture of the Chelungpu fault. It is an inclined borehole dipping 50° toward the west and the total drilling depth is 455.3 m (Fig. 6). Detailed core descriptions and drilling conditions have been documented by Tanaka et al. (2002), Hwang et al. (2002), and Heermance et al. (2003). Tanaka et al. (2002) considered that two rupture surfaces were possibly activated by the 1999 Chi-Chi earthquake, one at a 225-m drilling depth and the other at 330 m. The 225-m drilling depth rupture roughly coincides with the boundary of the Chinshui Shale and the Kweichulin Formation and the 330-m drilling depth locates at the Kweichulin Formation (Hwang et al. 2002). Tanaka et al. (2002) considered that the 225-m drilling depth coincides with a shallow seismic profile of high water content. Heermance et al. (Heermance et al. 2003) considered the depth around 330 m to be a high-density rupture zone that indicated earthquakes rupture.

In order to identify which rupture zone is associated with the Chi-Chi earthquake, we measured the depths of the earthquake rupture in three steps using the area balancing method. First, we projected the 330-m and 225-m drilling depths to the surface giving positions A and B, respectively. Position C was the location of the Fengyuan well (Fig. 6). Second, we made a vertical displacement profile from the drilling site to the surface rupture to calculate the increased area, which is parallel to the displacement direction. The GPS control point, L099, is near the Fengyuan site, and it shows a 7.72-m horizontal displacement in a 330° direction and a 3.3-m vertical displacement (LSB 2000). Thus, we use 7.72 m to be the horizontal displacement amount. According to scrap high and vertical displacement of L099 controlling point, we use 3 m, 3.3 m and 4 m to be the calculated values of the vertical displacement and assume that the vertical displacement is the same across the measured profiles. Thus, the increased area can be simplified as the distance from the surface rupture to the calculated positions multiplied by the vertical displacement. The measured depth of the earthquake rupture in each position is shown in Table 1 and Fig. 6.

The possible depths of the earthquake at position A are 241 m, 265 m and 321 m according to the 3-m, 3.3-m and 4-m vertical displacements. Figure 6 shows possible depth of the slip plane. It shows that the position of an earthquake rupture at 252.79 m inferred from Heermance et al. (2003) is located in our estimated depth but the position at 172.36 m inferred from Tanaka et al. (2002) being unable to fit into our estimated depth. Using the same concept, we can estimate that if the depth at 172.36 m is the slip plane, then the horizontal slip amount would be 15.2 m according to a 3.3-m vertical displacement or the vertical displacement would

Vertical	Length of measuring profile		
displacement	A:619 m	B:795m	C:1038m
3m	241m	309m	403m
3.3m	265m	340m	444m
4m	321m	412m	538m

Table 1. Depths of the earthquake rupture according different vertical displacements at different positions. A, B, and C are the measured sections.

be about 6.5 m based on a 7.22 m horizontal displacement. These values are not consistent with coseismic displacement of GPS. The result prefers the depth at 252 m (330-m drilling depth) as being the possible slip plane, which is similar with a depth of 265 m obtained by assuming the 3.3-m vertical displacement. We thus use the 3.3-m vertical displacement to estimate the geometry of the fault plane. The dip angle is about 37° from position C to B and near 50° from position B to the surface rupture. Another Chelungpu fault drilling (BH2) is located only 40 m to the west of the surface rupture and this well penetrates the fault zone at a 50-m depth indicating that the dip angle is 50° (Lee 2000a) (Fig. 7). This dipping angle coincides with our results confirming the usefulness of the area balancing method in this case.

Tanada et al. (2002) inferred from the shallow seismic profile that the 225-m drilling depth was the position of the earthquake rupture. The problem here being that the drilling direction is in a westward direction, but the seismic profile is in a near 297° direction (Fig. 8). Therefore, we cannot connect both data directly. In order to accommodate both, we would have to either project the well data in a 297° direction or project the seismic data in a westward direction. In order to compare our results with the seismic profile, we also measured the depth of the slip plane along the seismic profile by the area balancing method. The black dash line is the fault plane according to the 3.3-m vertical displacement. From a depth of 180 m to one of 444 m, it shows bedding slip; and from 180 m to the surface it shows about a 50° dip angle (Fig. 8).

6. CONCLUSION AND DISCUSSION

According to the area balancing method, we suggest that the 330-m drilling depth at Fengyuan well is the position of the Chi-Chi earthquake rupture.

The major problems with using an area balancing method are the accuracy of the vertical and horizontal displacements from the surface to the slip plane. In this paper, we assume that the horizontal displacement is the same from surface to slip plane at depth and the same from the surface rupture to measuring points. The GPS data show that the horizontal displacement decreases progressively from the surface rupture toward the east (Yu et al. 2001) and the changing rate of horizontal displacement parallel to the displacement direction is about

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Fig. 7. (a) The wells position of BH1 and BH2. (b) The vertical displacement is about 3 - 4 m along the BH1 section and the fault dip angle is near 50° (Lee et al. 2000).

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Fig. 8. (a) The position of the shallow seismic profile around Fengyuan site (Wang et al. 2002). (b) The seismic profile is according to Wang et al. 2002. The geometry of fault plane according to the area balancing method with 3.3-m vertical displacement. The black dash line shows the possible fault trace of the Chi-Chi earthquake. The gray color lines show the structures interpreted by Wang et al. (2002).

 $10^{-3} \sim 10^{-4}$ indicating that while the measuring profile is short, we can assume the horizontal displacement is the same; but if we want to measure a long profile, we should consider changes in the horizontal displacement.

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